

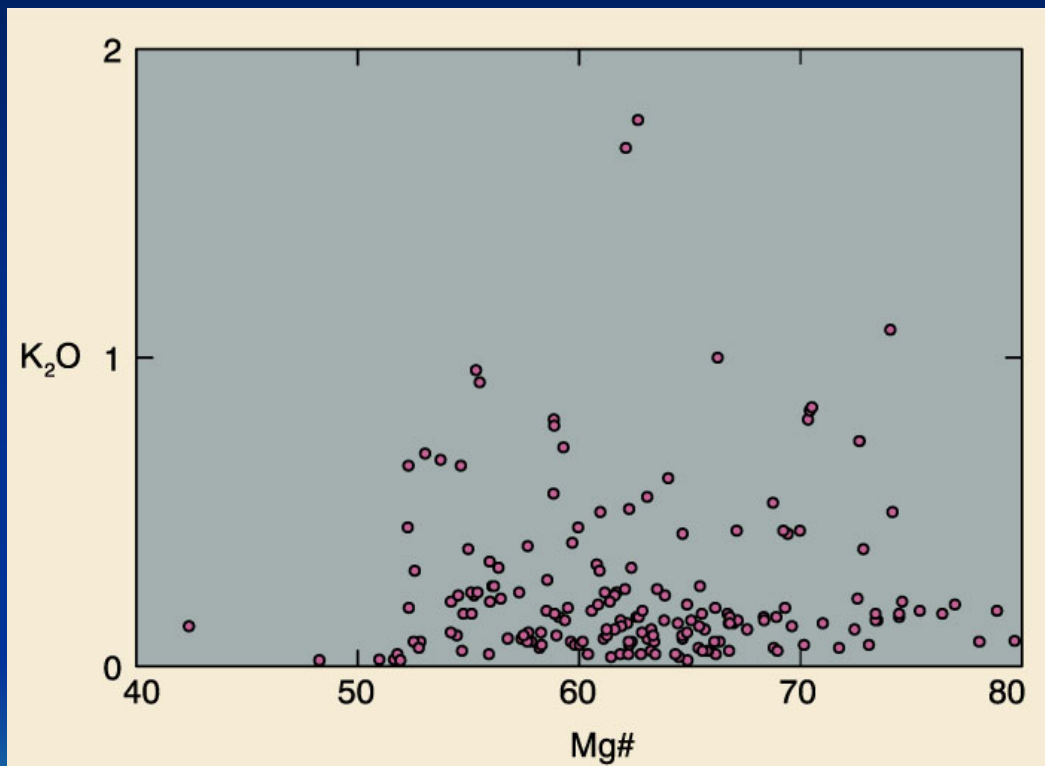
Geochemistry of MORBs

For constant Mg# considerable variation in K_2O

K: incompatible element

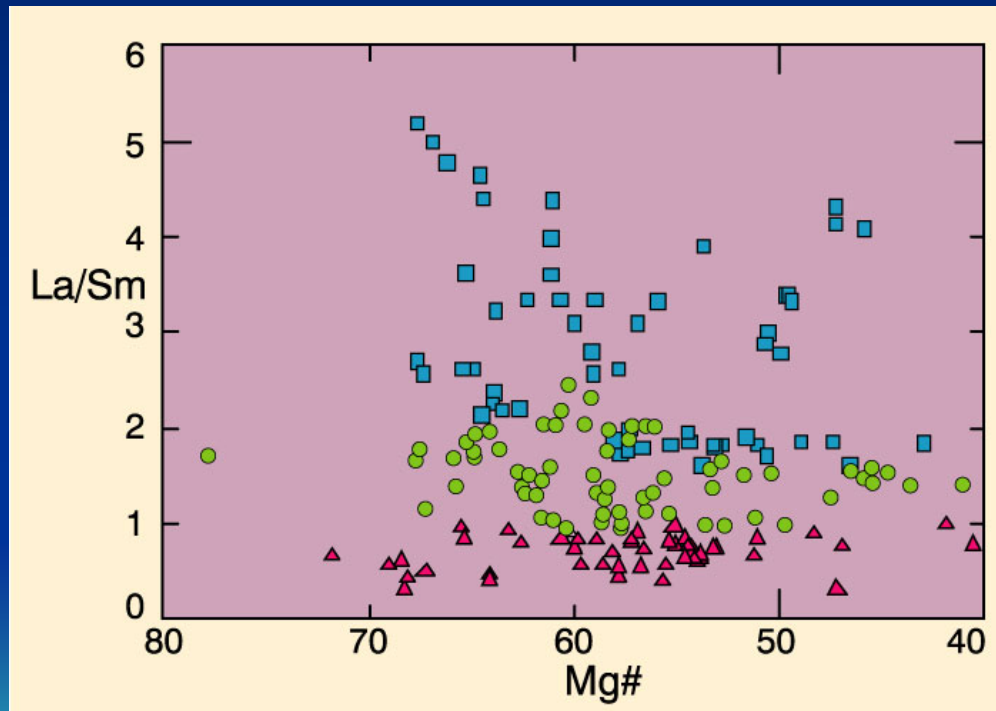
K concentration not affected greatly by <50% fractional crystallization

→ K variation reflects characteristics inherited from the mantle source



Trace Elements

E-MORBs (squares) enriched over N-MORBs (red triangles) and T-MORBs (green dots) regardless of Mg#



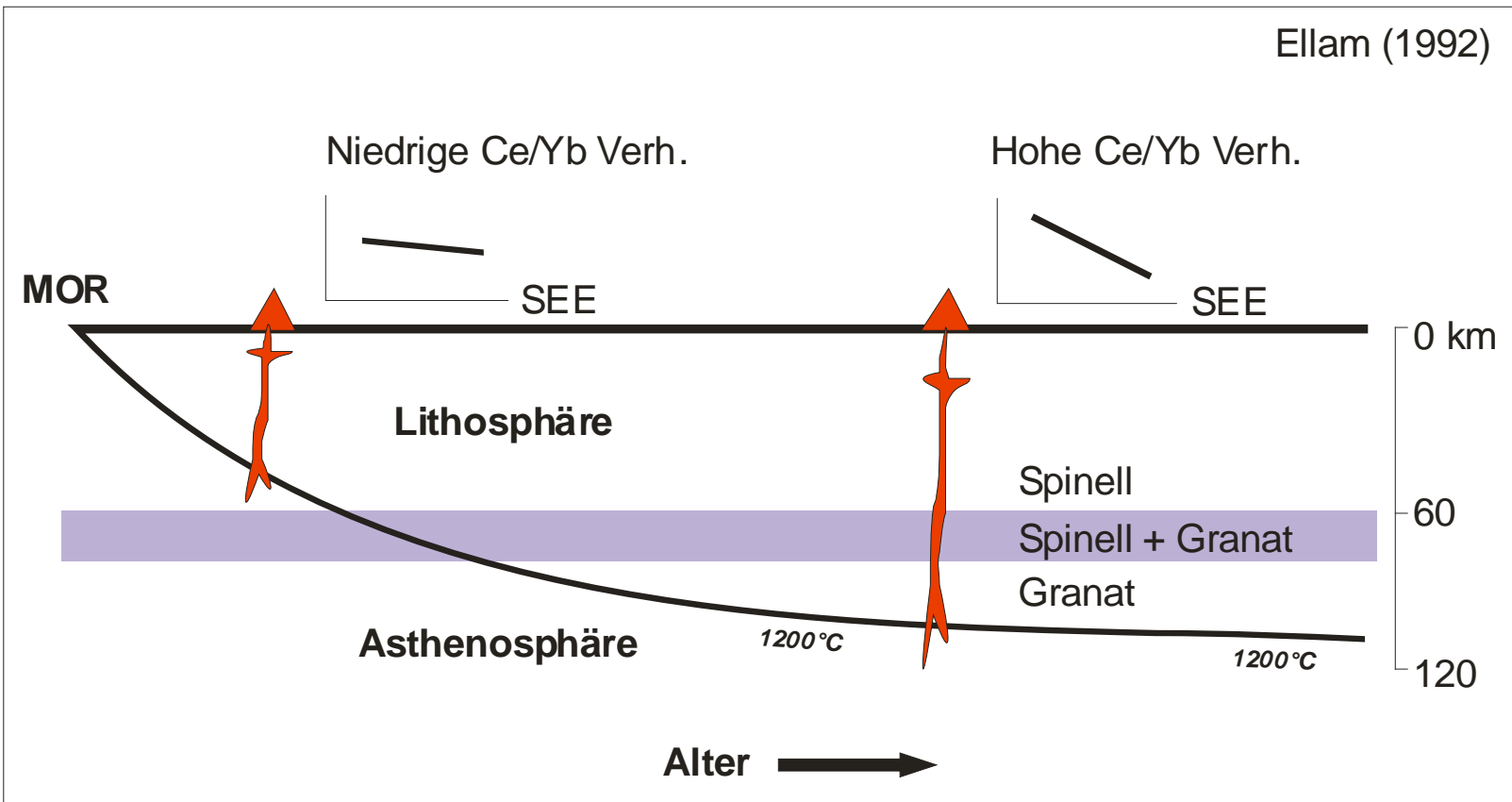
E-MORBs $\text{La/Sm} > 1.8$

N-MORBs $\text{La/Sm} < 0.7$

T-MORBs (transitional)
intermediate values

Data from Schilling et al. (1983)

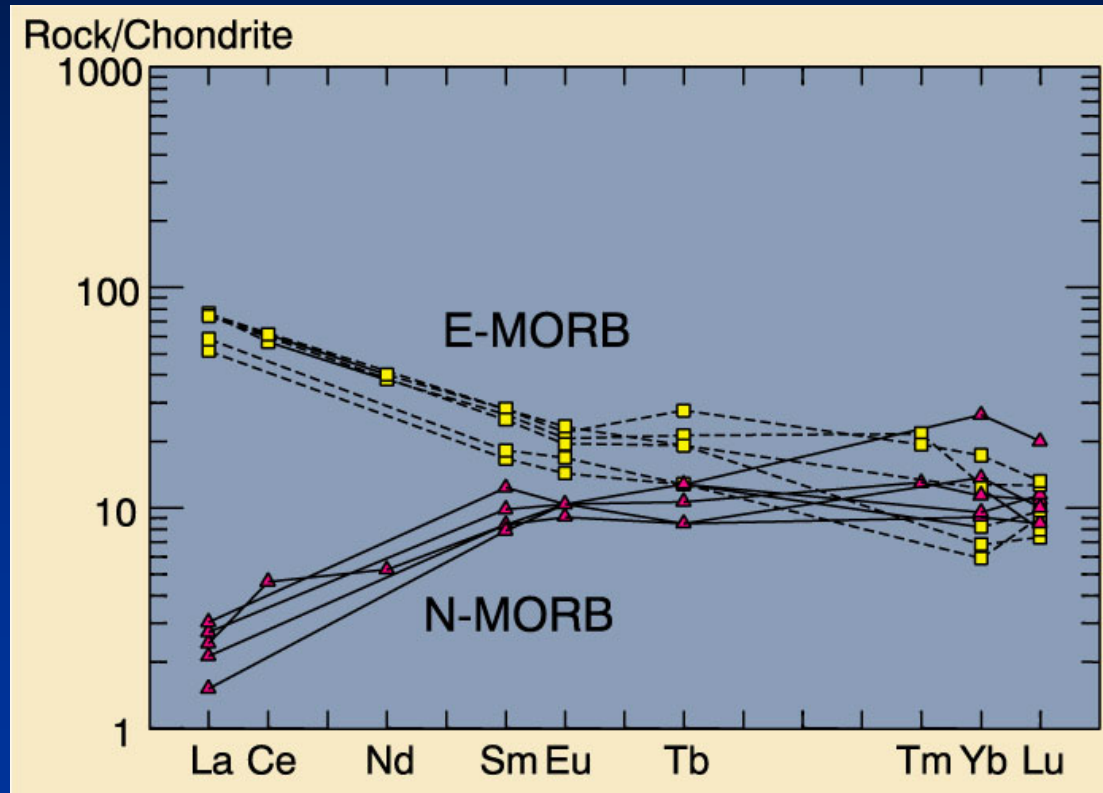
Trace element ratios and depth of melt generation



HREE depletion expected if garnet was a residual phase in the garnet-lherzolite field
Eu-anomaly expected if plagioclase was residual in the plagioclase-lherzolite field

Trace elements “Schilling-effect”

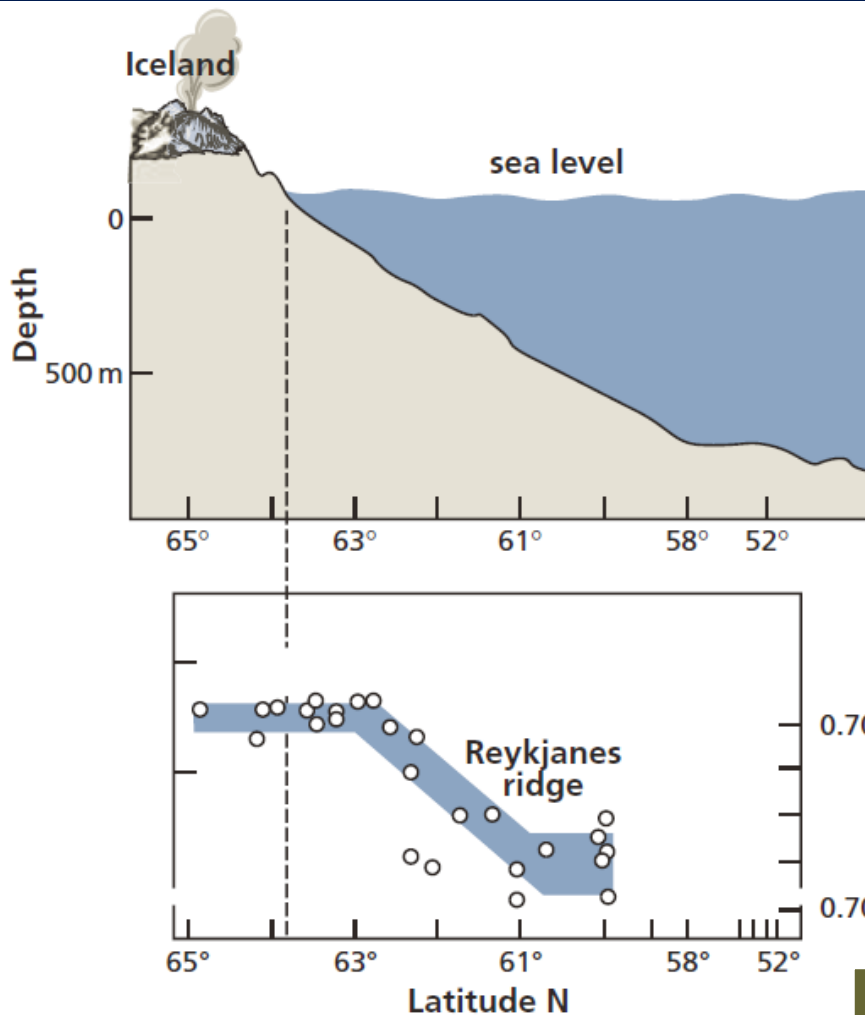
REE diagram for MORBs



Data from
Schilling et al.
(1983)

Origin of MORB in the **Spinel lherzolite field** is compatible with the REEs.

E-MORB or “Schilling-effect”



Incompatible-rich and incompatible-poor mantle source regions for MORB magmas

textbook
J.C. Allègre

Conclusions about MORBs, and the processes beneath mid-ocean ridges

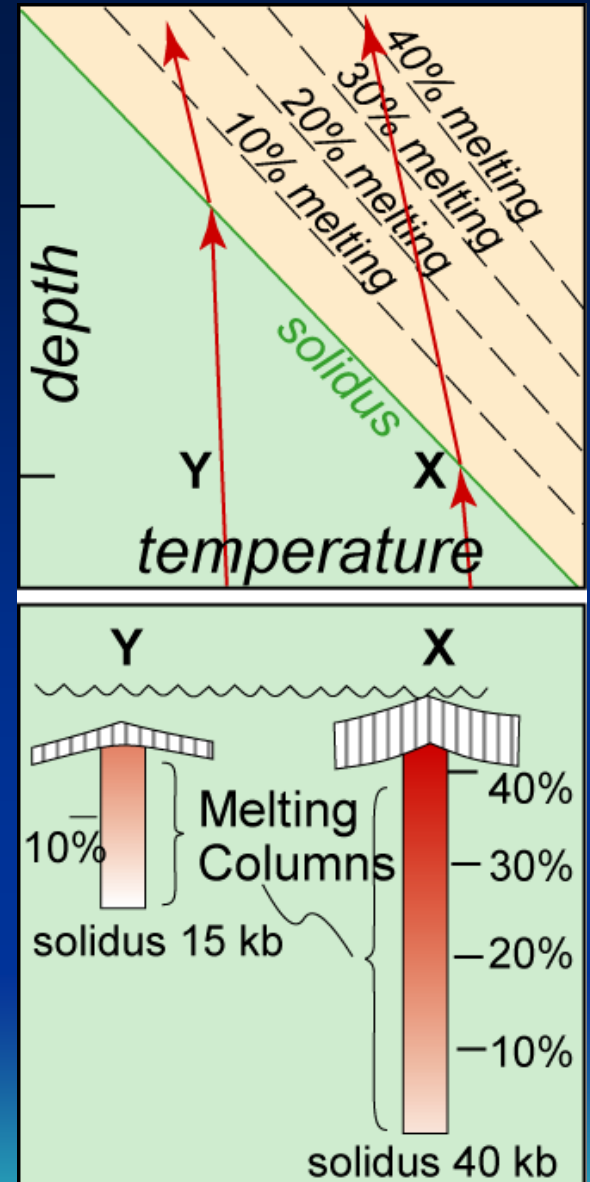
The mantle beneath the ocean basins is not homogeneous

- N-MORBs tap an upper, depleted mantle
- E-MORBs tap a deeper enriched source
- T-MORBs = mixing of N- and E- magmas during ascent and/or in shallow chambers



Mantle melting in reality

1. Melting is polybaric, i.e. occurs at a range of depths.
2. Melt can react with residual solid during transport
3. K_D 's may vary with depth due to changes in P, T, and composition of residual phases.
4. Phases can become exhausted during progressive melting



REE variations during melting

Modeled REE variations during different degrees of melting of a garnet peridotite mantle source

	wt%
olivine	60
opx	20
cpx	10
garnet	12
total	100

